

Aerosols, CCN and fog, oh my!

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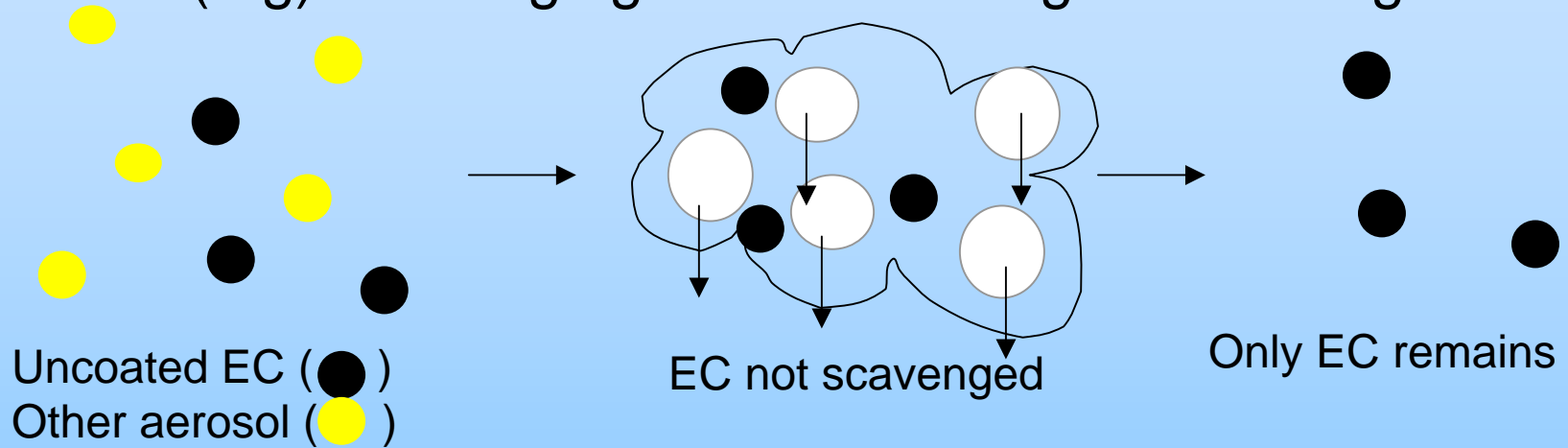
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Aerodyne & UMIST
Ells Dutton
Andreas Stohl

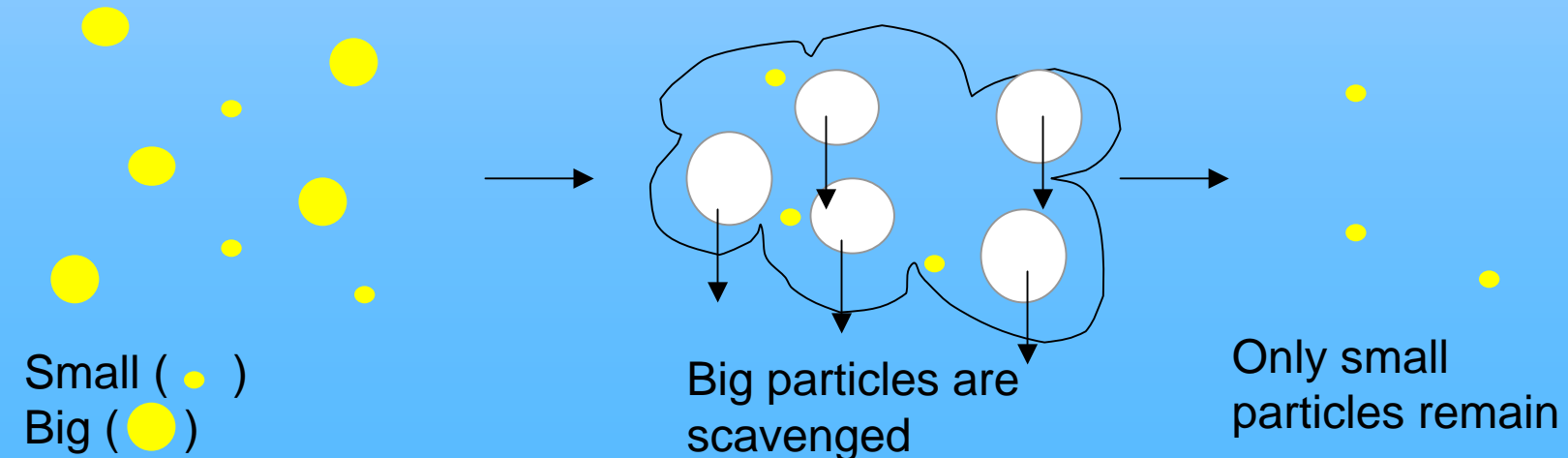


Hypotheses

Cloud (fog) scavenging decreases single scattering albedo.

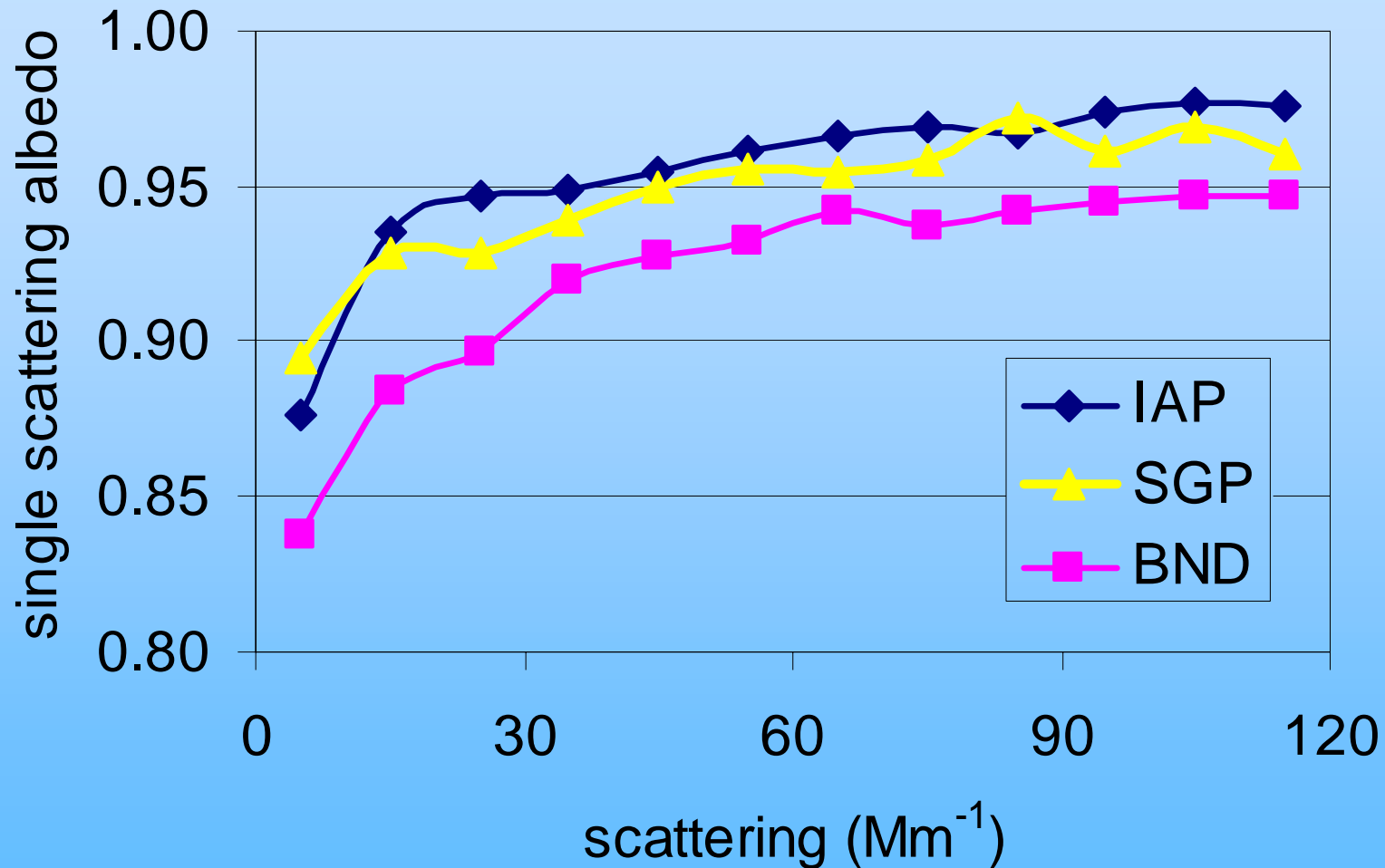


Cloud (fog) scavenging increases backscattering fraction.



Single-scattering Albedo

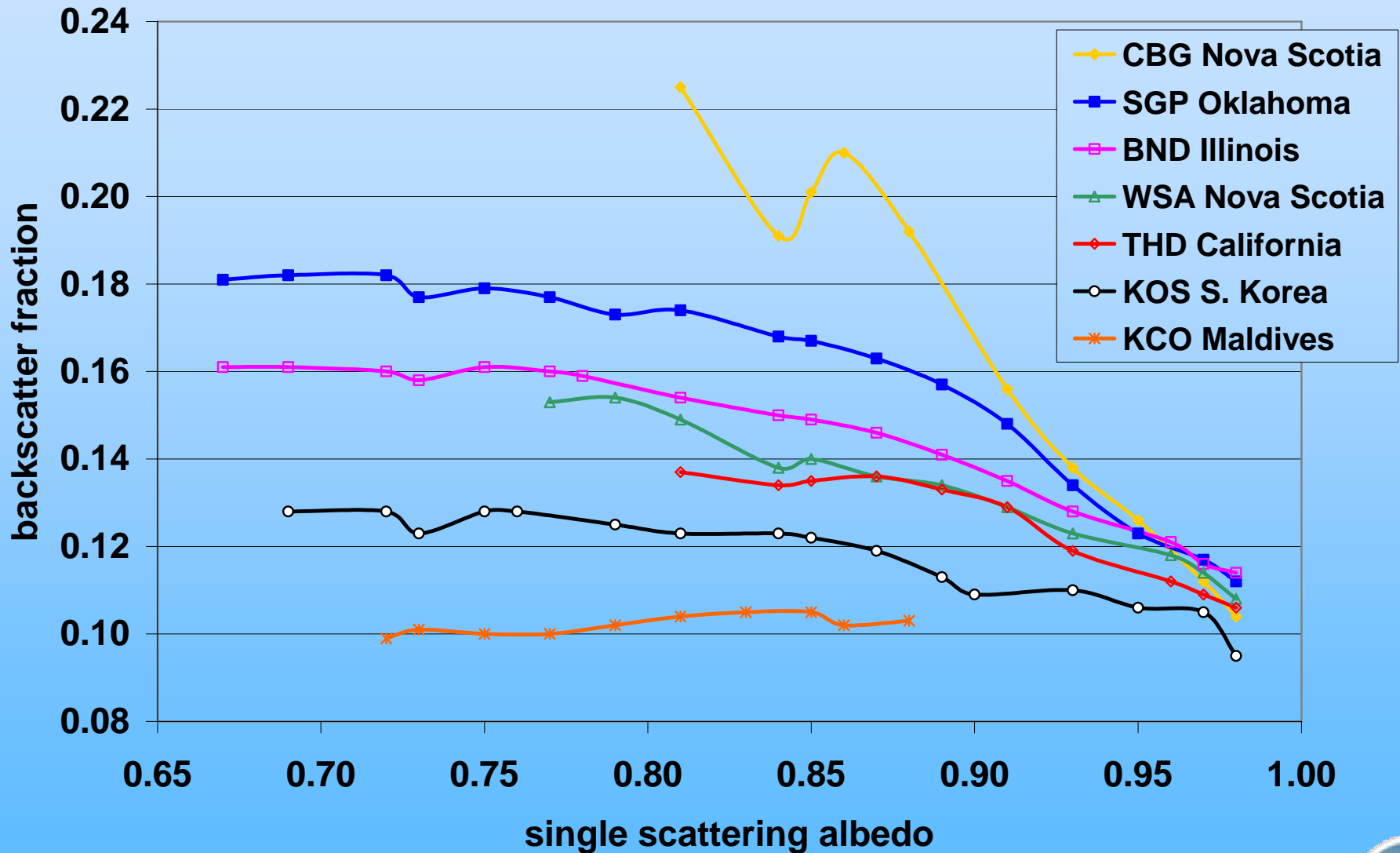
Increases as Pollution Increases



Notes: Binned, daily mean results for RH<40%, D_p<1μm, and λ=550nm.
BND=Bondville (IL), 1994-2003; SGP=Lamont (OK), 1996-2003;
IAP=Lamont (OK) vertical profiles, 2000-2004.



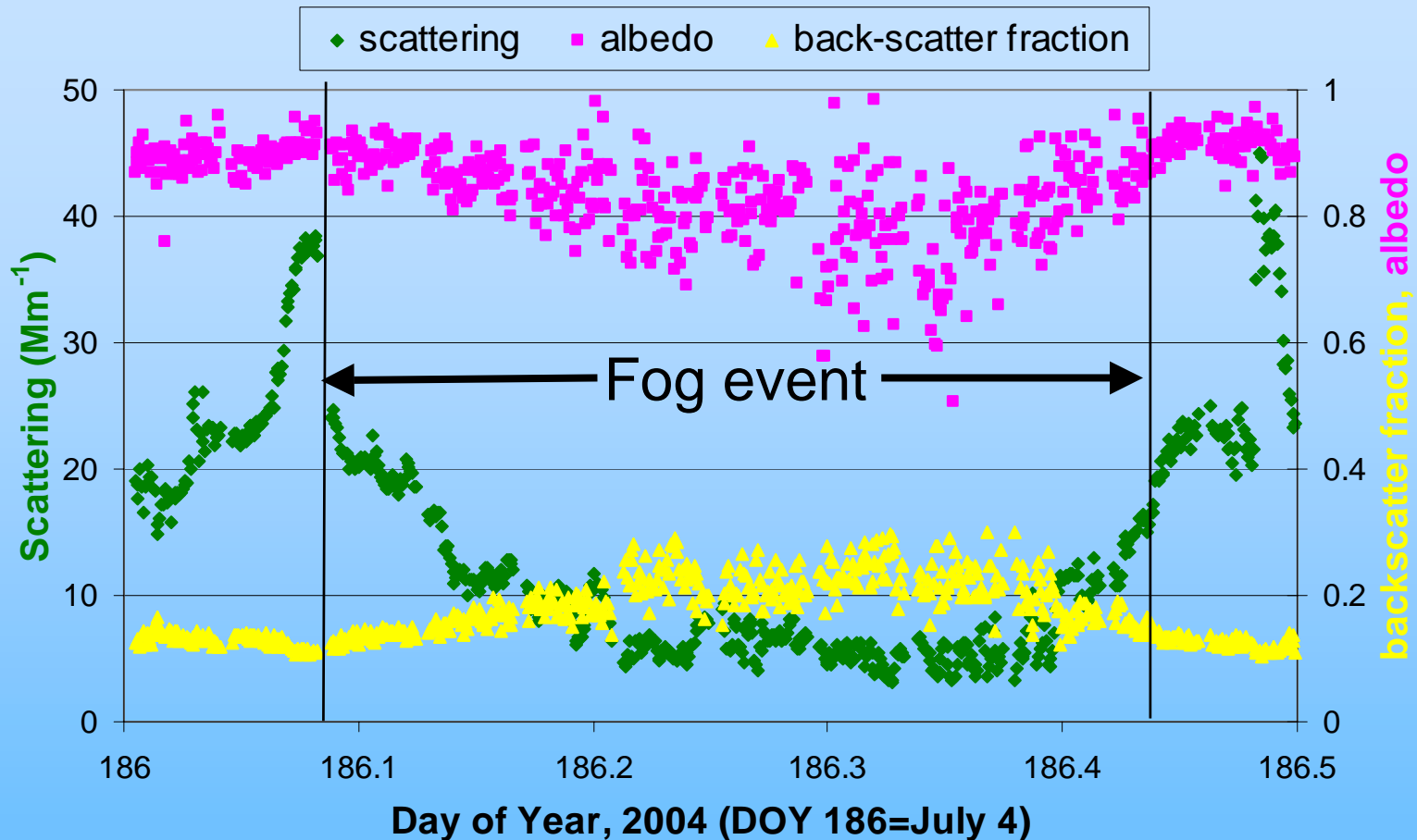
Co-variance of Aerosol Radiative Properties



Values measured at 550 nm wavelength, $RH < 40\%$, $D_p < 10 \mu m$



Effect of fog on aerosol optical properties

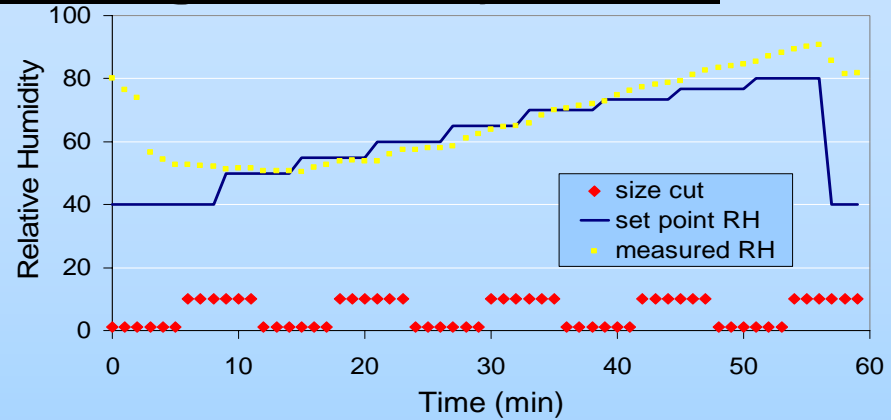
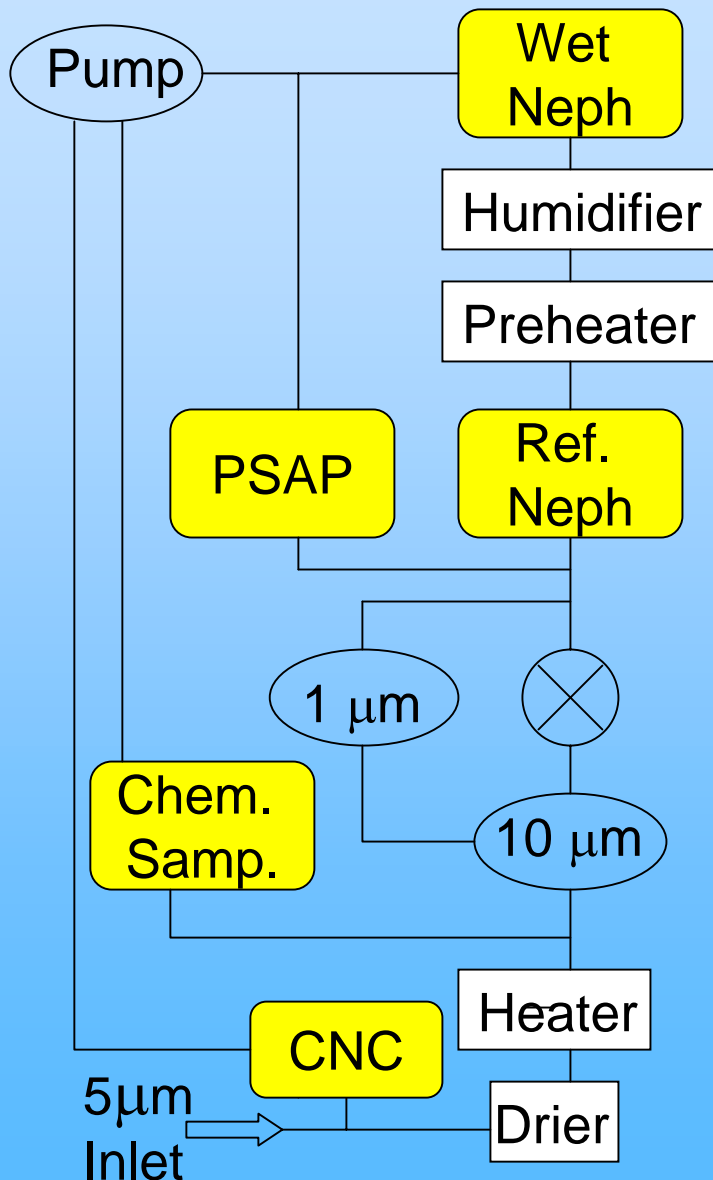


Onset of fog causes:

- Decrease in light scattering,
- Increase in back-scatter fraction,
- Decrease in single scattering albedo



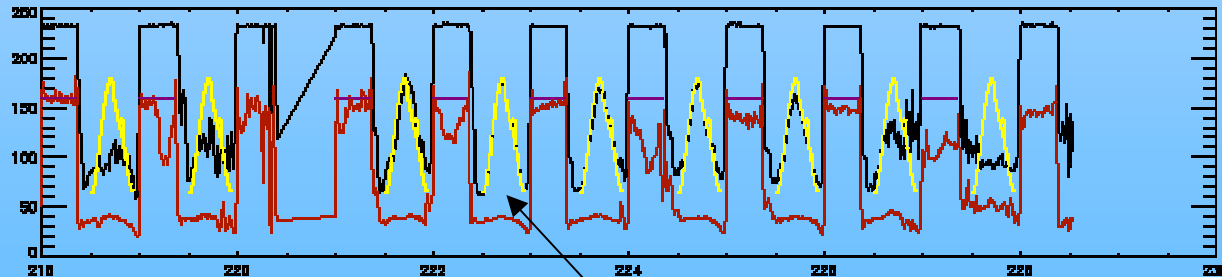
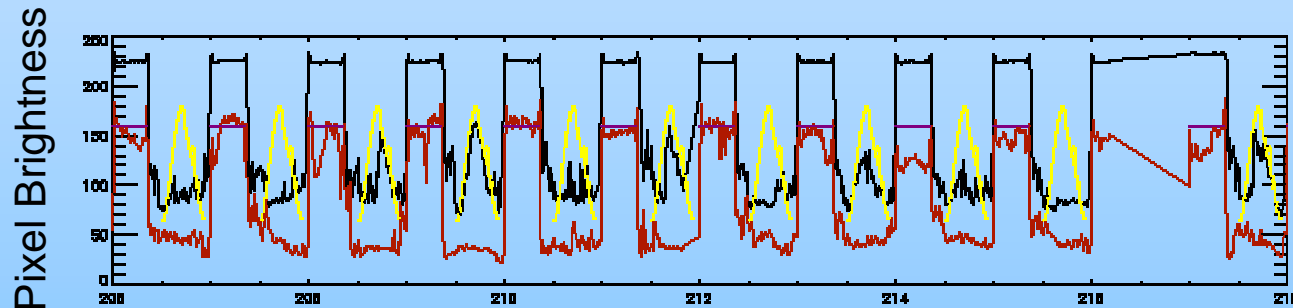
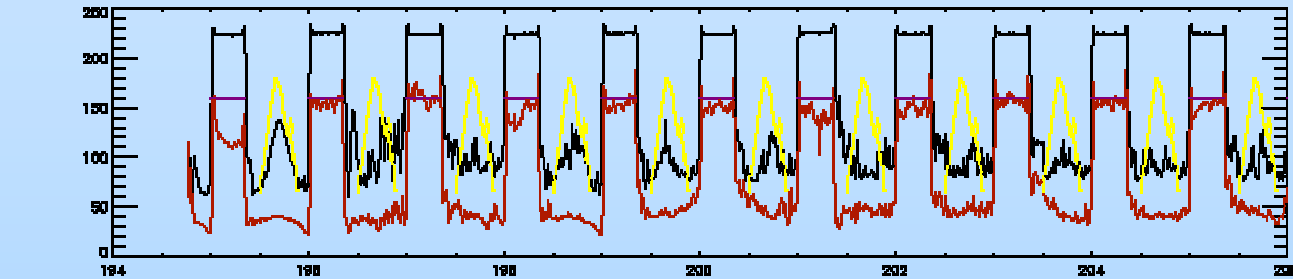
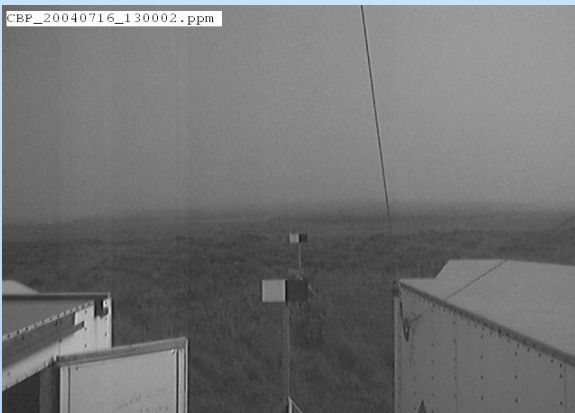
CMDL Humidograph System



Plus...
 →CCN
 →SEMS
 →FogCam



Fog at Chebogue point



Clear day

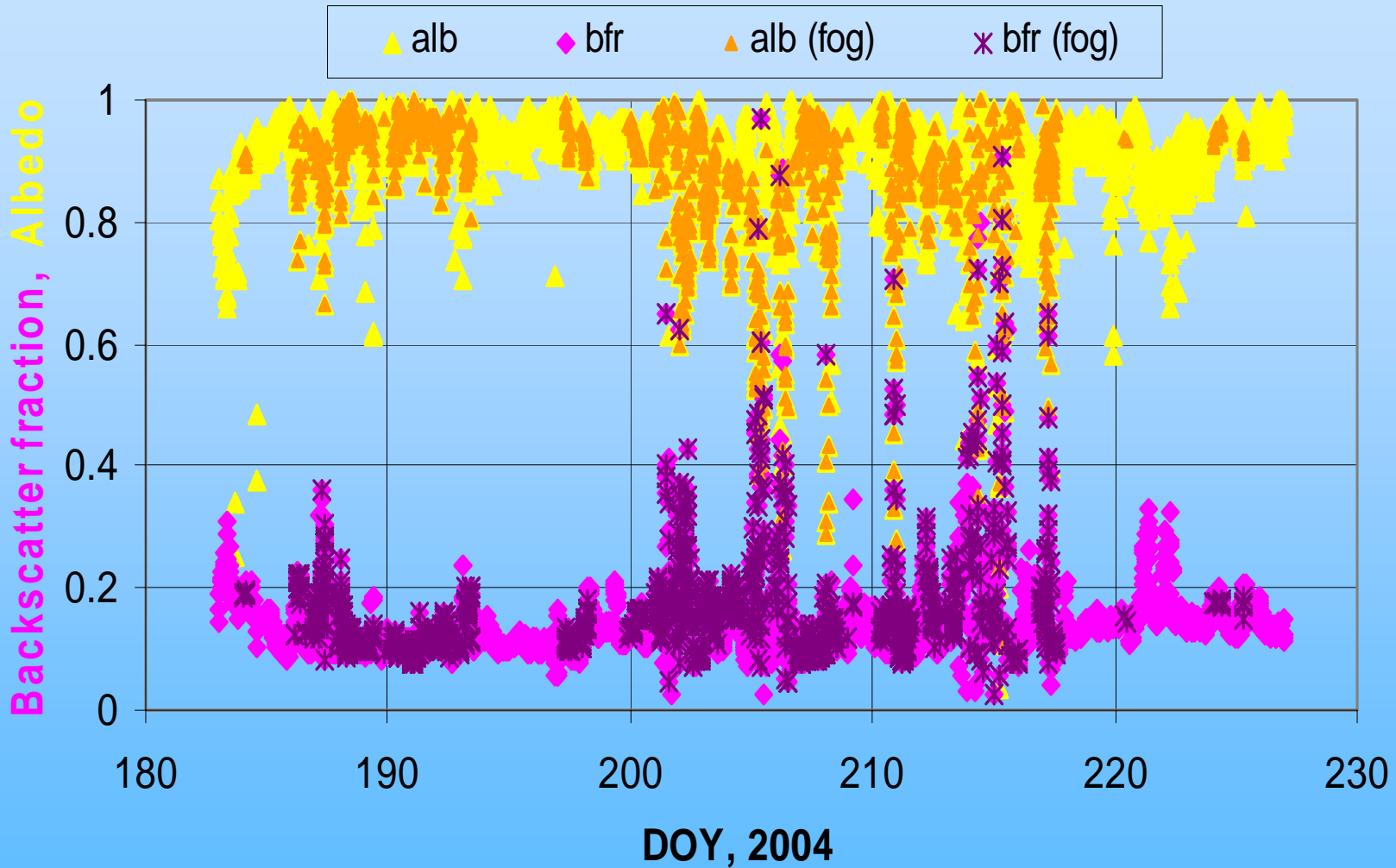
Webcam –
2 black/white targets
1 picture/minute
Nighttime illumination



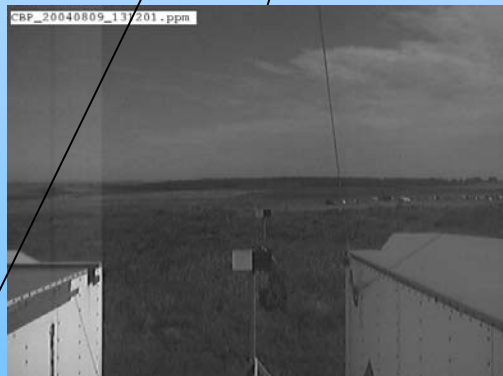
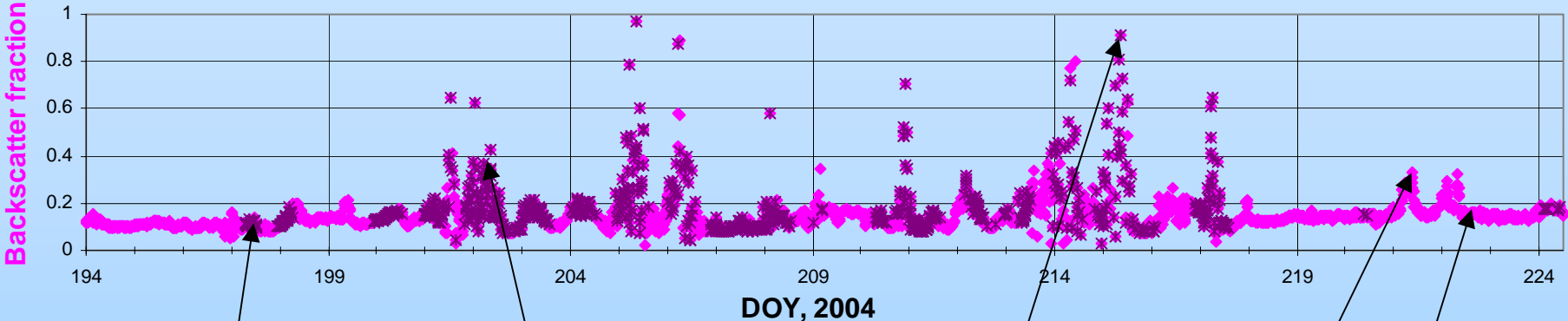
We will use Ellis Dutton's fog identification based on IR fluxes instead of webcam.



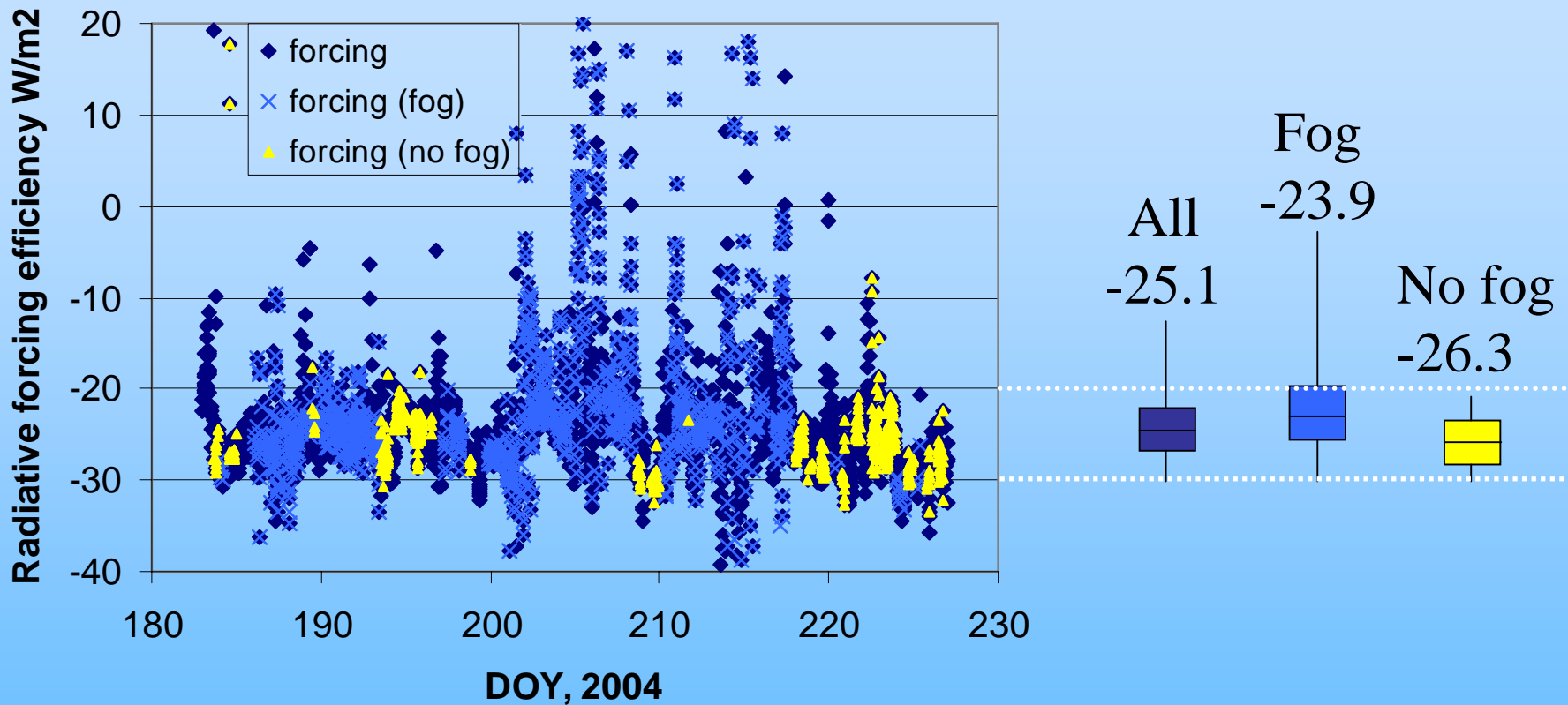
Effect of fog on aerosol optical properties



Webcam versus IR flux



Radiative Forcing at Chebogue Point

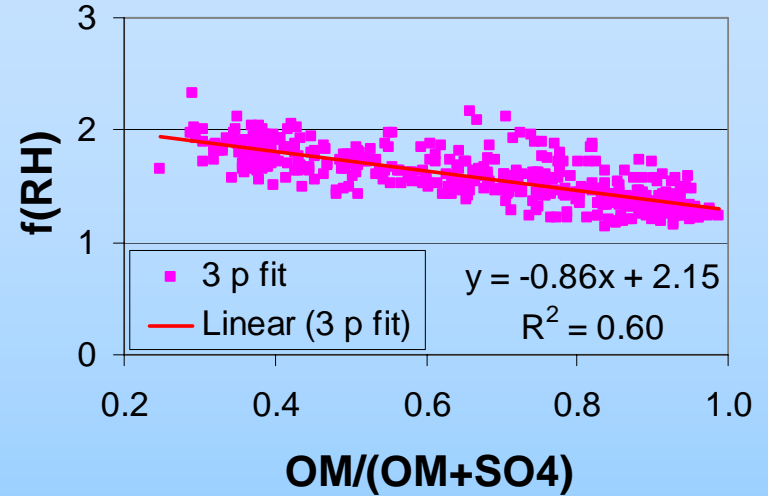
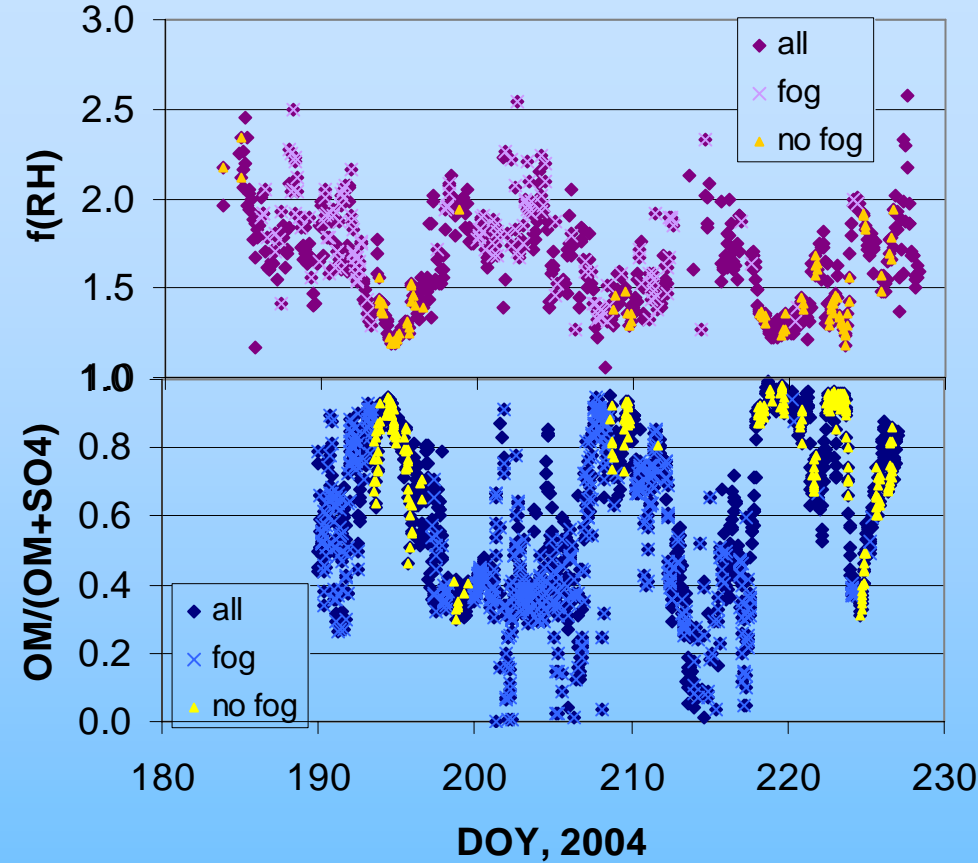


$$\Delta F \approx -DS_0 T_{at}^2 (1 - A_c) (1 - R_s)^2 \tilde{\omega}_0 \bar{\beta} \delta \left[1 - \frac{2R_s}{(1 - R_s)^2} \left(\frac{1 - \tilde{\omega}_0}{\tilde{\omega}_0 \bar{\beta}} \right) \right]$$

From Haywood and Shine, 1995



Fog, chemistry and hygroscopic growth



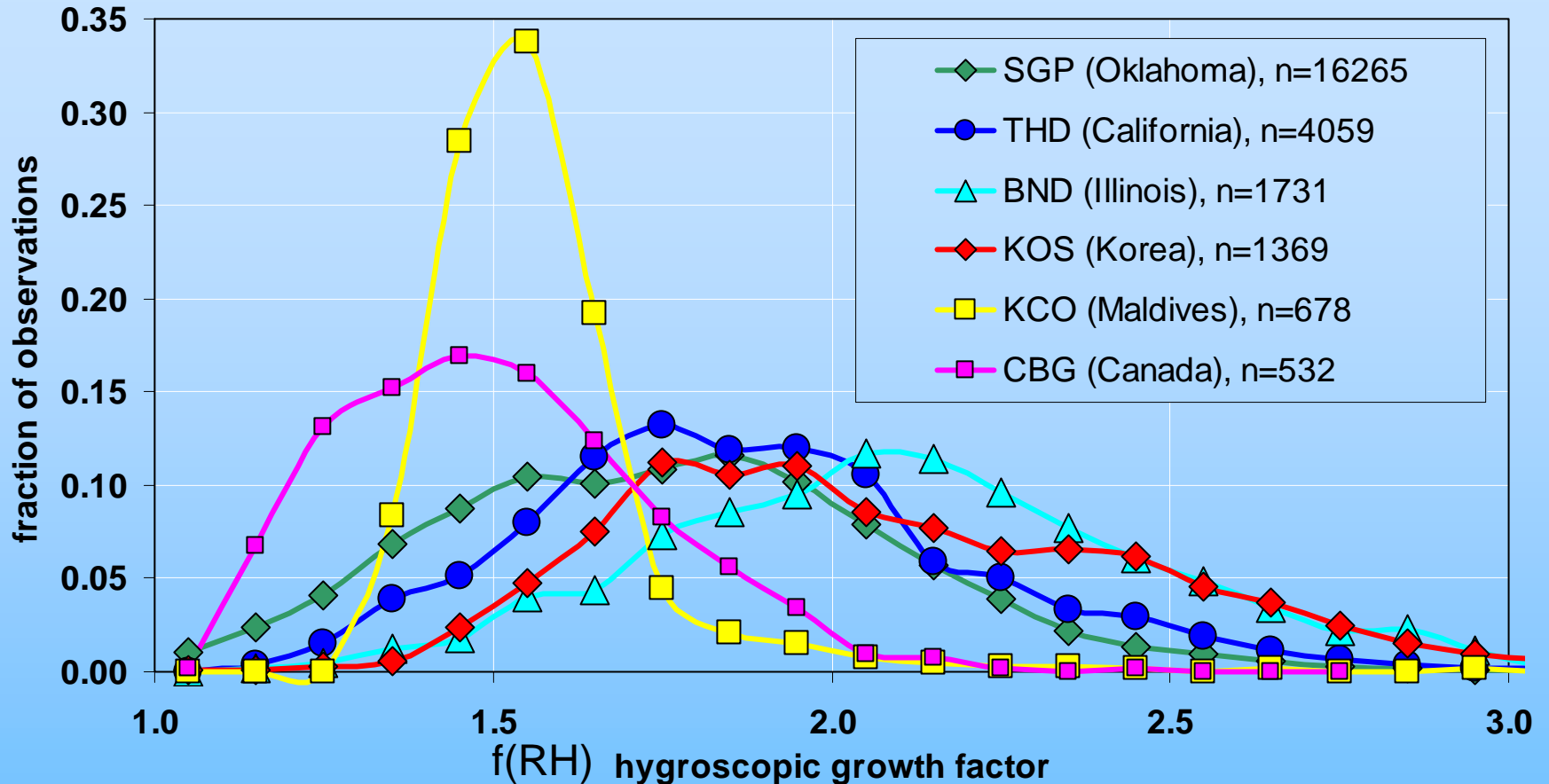
- $f(\text{RH})$ correlated with $\text{OM}/(\text{OM}+\text{SO}_4)$ ratio
- Highest $\text{OM}/(\text{OM}+\text{SO}_4)$ ratio during non-foggy periods
- $f(\text{RH})$ higher during foggy periods!

SO_4 & OM data from Aerodyne/UMIST
[OM] > 0.25 $\mu\text{g}/\text{m}^3$, $f(\text{RH})$ for $d_p < 1 \mu\text{m}$

Another hypothesis – MSE curve?



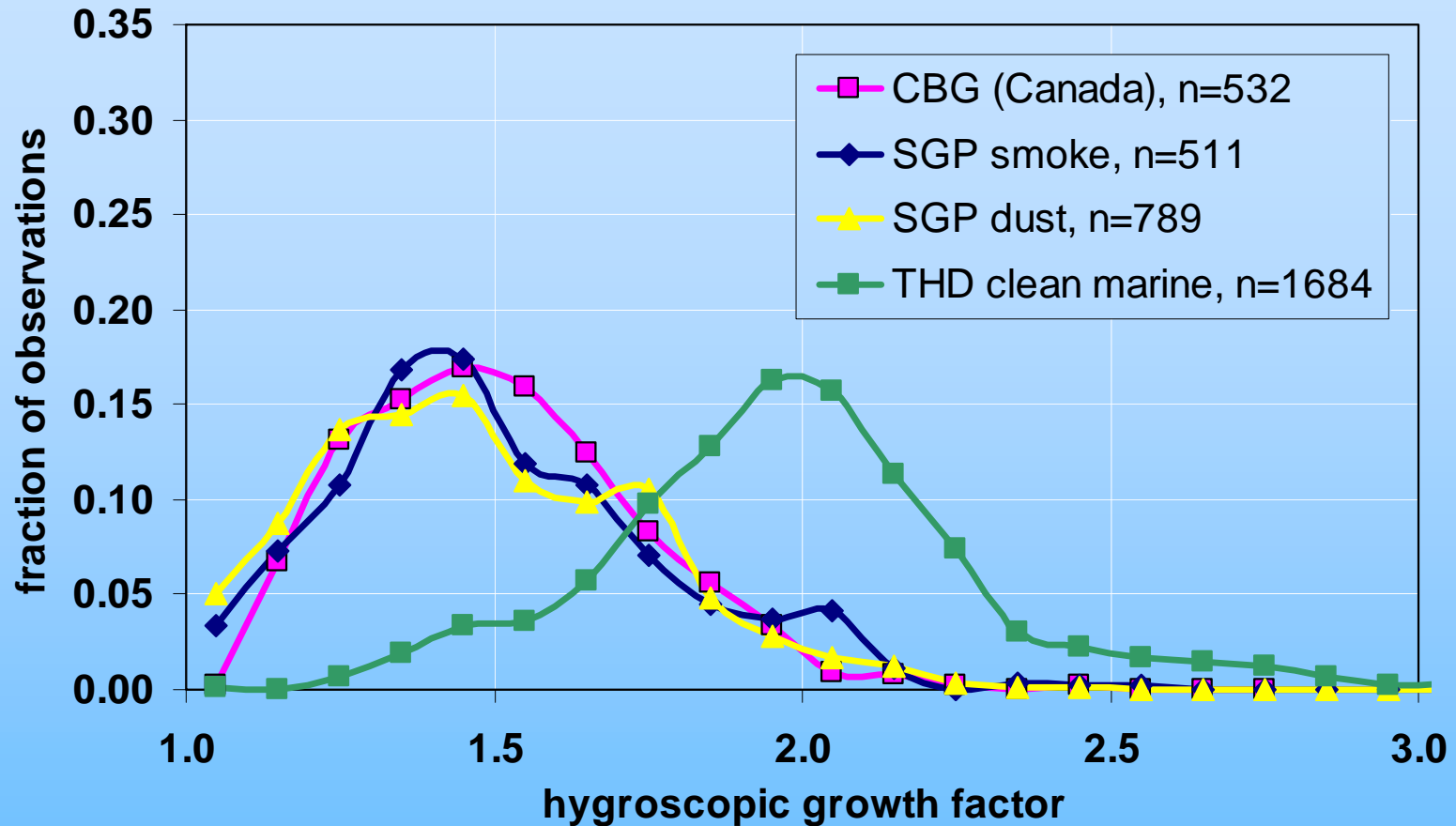
Regional Differences in Hygroscopic Growth



Parameter plotted is $\sigma_{sp}(85\% \text{ RH})/\sigma_{sp}(40\% \text{ RH})$, based on 2 parameter fit.
BND data courtesy of Prof. Mark Rood, Univ. of Illinois



Hygroscopicity as a function of aerosol type



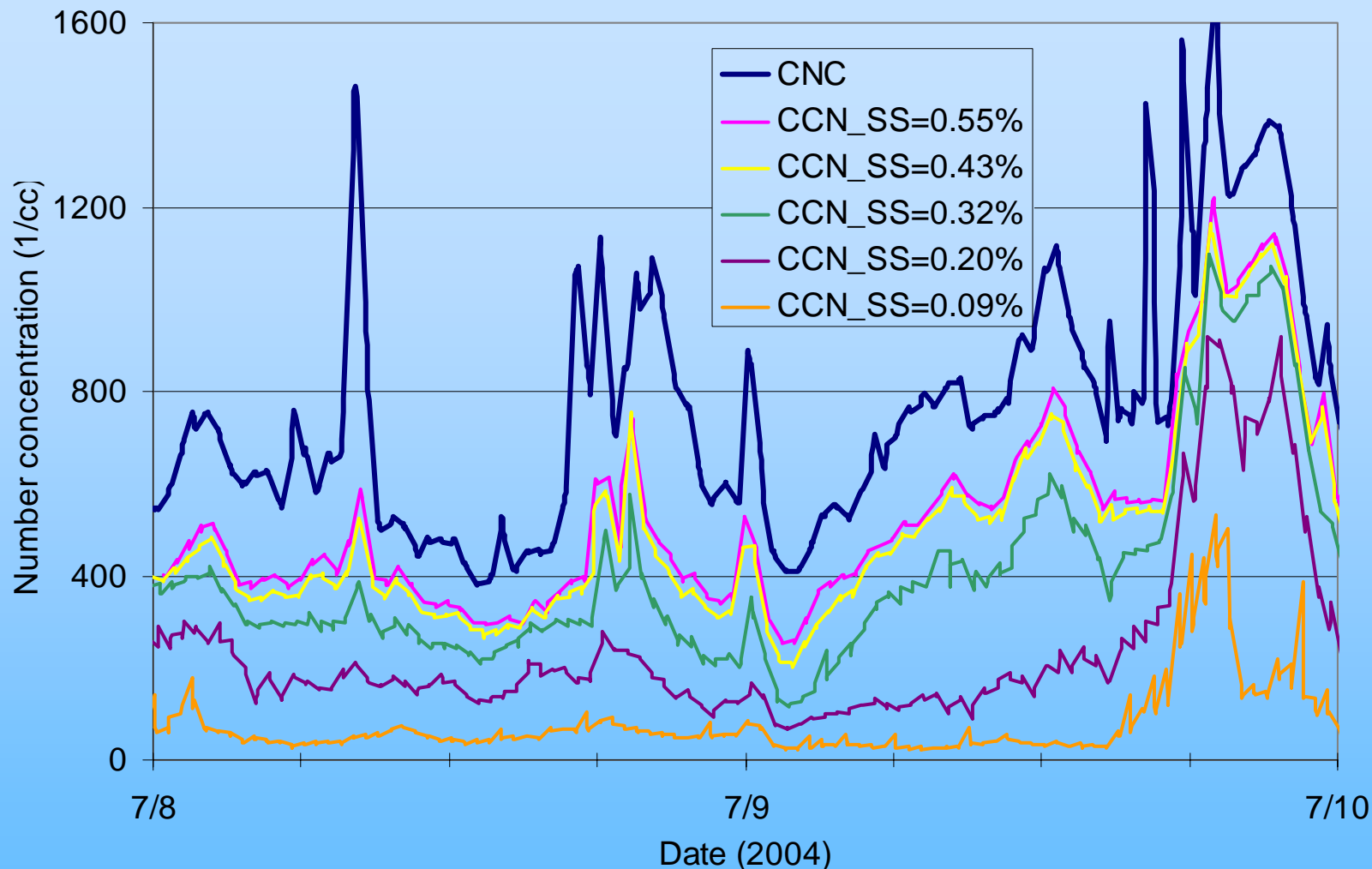
Smoke: single scattering albedo < 0.92

Dust: sub-micron scattering fraction < 0.65

Clean marine: single scattering albedo > 0.96 , sub-micron scattering fraction < 0.65



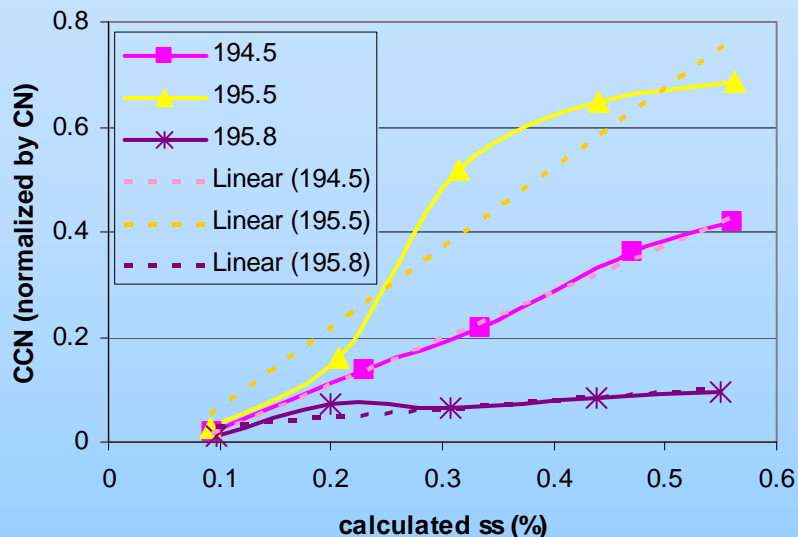
CN and CCN time series



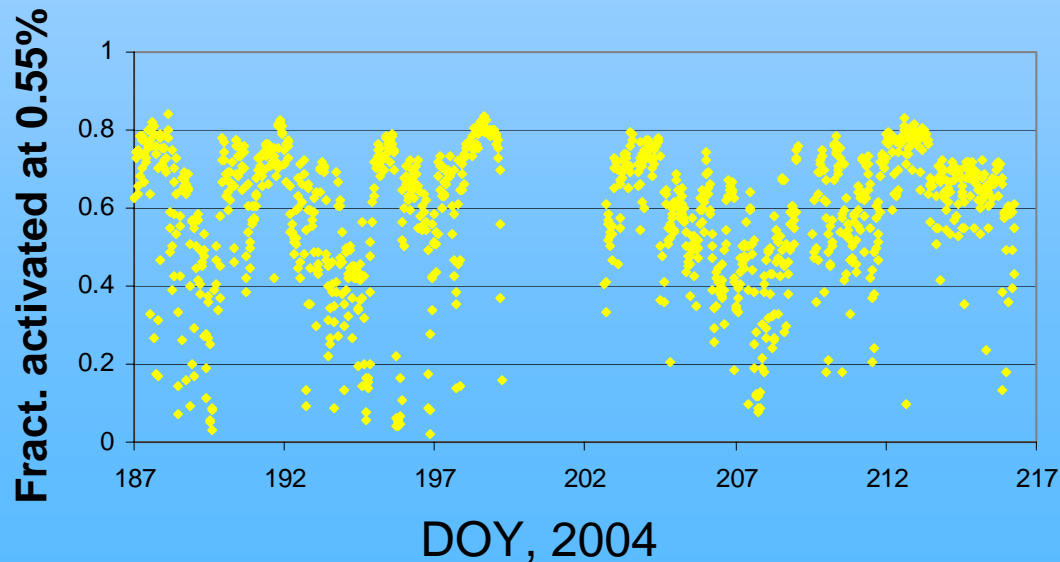
- Condensation nuclei concentration (CNC) provides upper envelope to CCN concentrations



CCN spectra at Chebogue Point



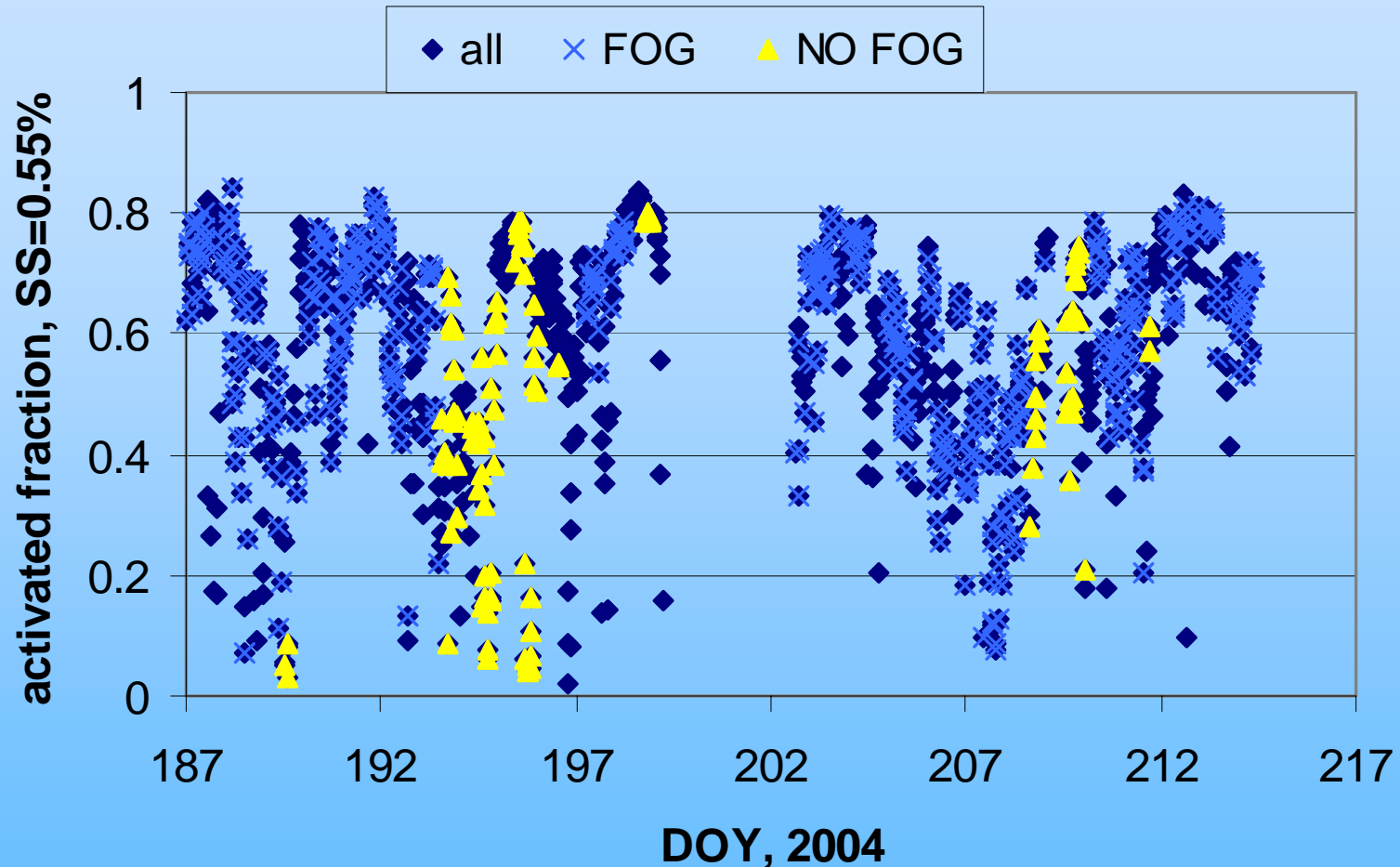
- Normalized CCN by CN
- Fit straight line as first approximation
- Lots of variation in slope



QUESTION:
How do spectra change during fog/aerosol events?



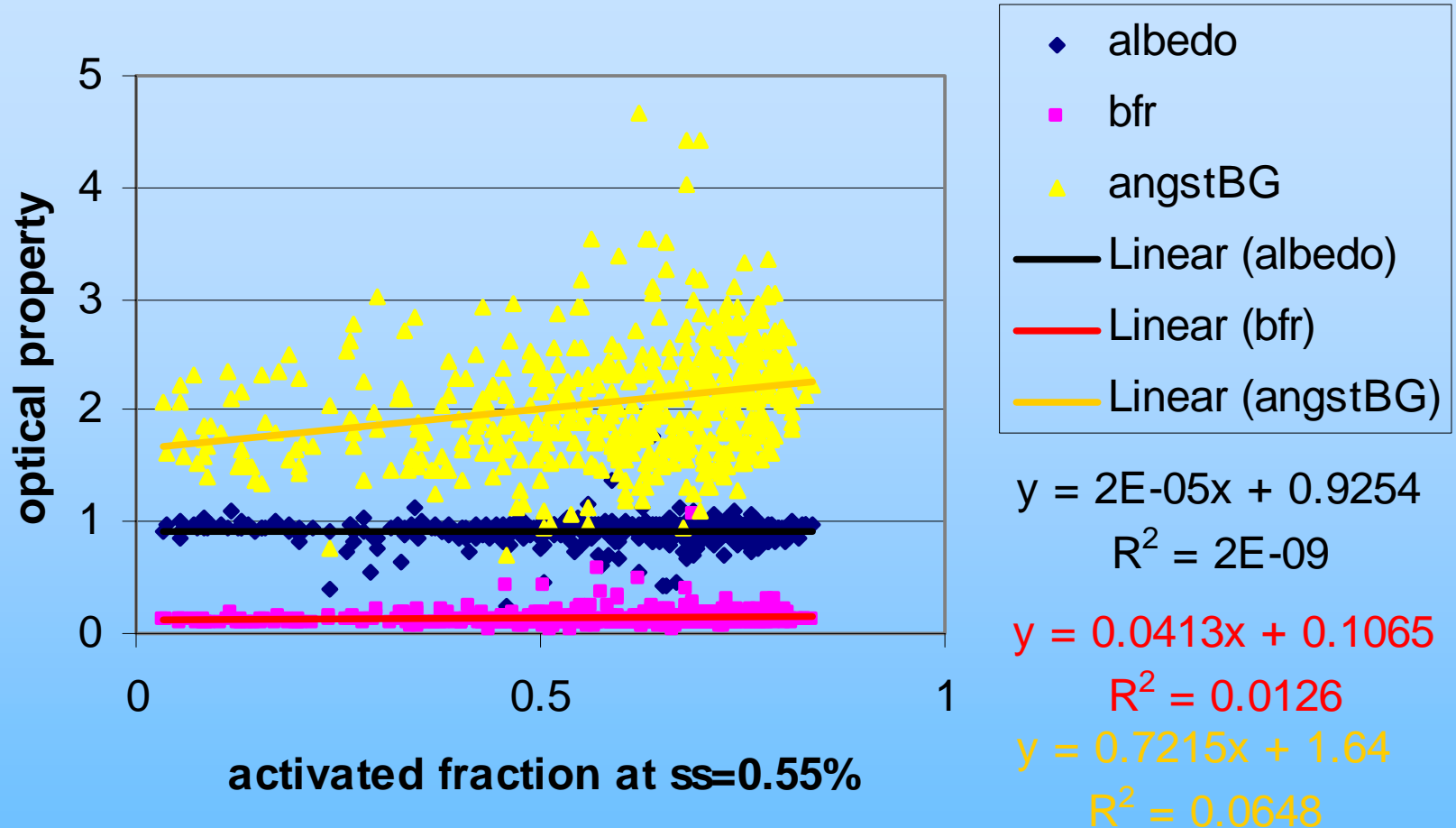
Activated fraction at SS=0.55% and fog



There doesn't appear to be a relationship between the fraction of CCN that activate at 0.55% super-saturation and the presence or absence of fog.



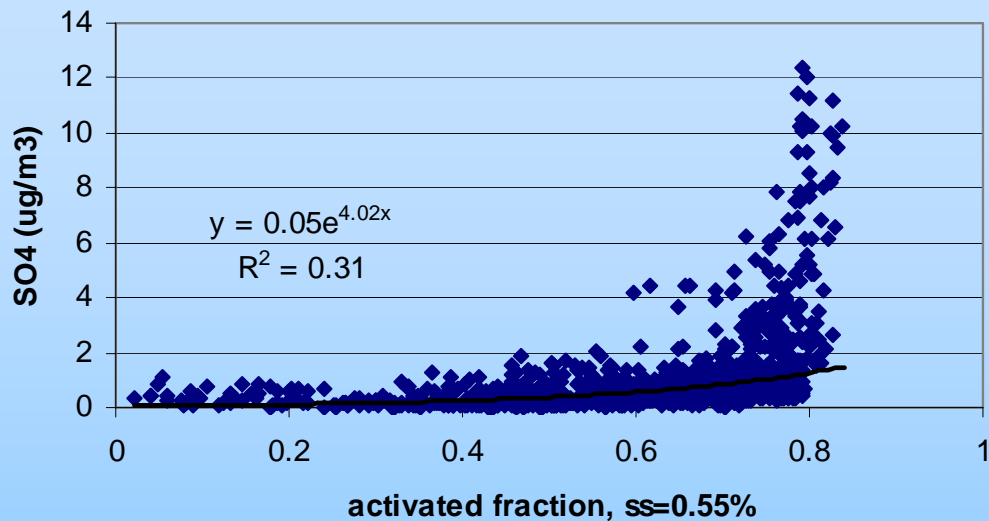
Activated fraction and aerosol optical properties



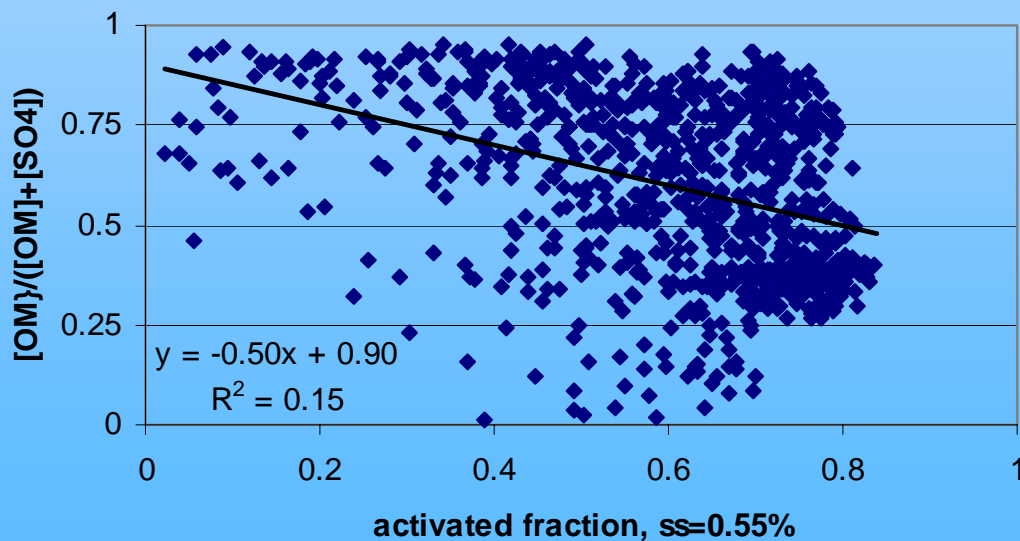
The **Ångström exponent** is inversely related to particle size. The activated fraction appears to increase slightly when the aerosol population is dominated by smaller particles.



Activated fraction and aerosol chemistry



More sulfate = More CCN



Higher
 $\frac{[OC]}{([OC]+[SO4])}$ = Less CCN



Summary

- Fog has a significant influence on aerosol optical properties
- Based on simplistic model, does not appear to strongly affect radiative forcing efficiency
- Hygroscopic growth is strongly correlated with aerosol composition – OM reduces hygroscopic growth
- CCN spectra are variable; but at this stage we're not sure what the sources of variability are, aside from general indications of size and chemistry effects – Jose Jimenez's group (CU) will be exploring this.



THANKS!



Cloud Scavenging Decreases Single-scattering Albedo

- **If the particles are externally mixed**
 - aerosol light scattering is dominated by particles that are readily-scavenged by clouds, such as sulfates and water-soluble organics
 - aerosol light absorption is dominated by less readily-scavenged particles, such as graphitic carbon (soot)
- **If the particles are internally-mixed**
 - larger particles are scavenged more efficiently than smaller particles
 - the scavenged particles have much higher single-scattering albedos than the smaller particles
- **Cloud droplets are therefore enriched in light scattering particles relative to light absorbing particles**
- **When precipitation falls, it removes more of the light scattering particles than the light absorbing ones**
- **Cloud scavenging therefore systematically decreases aerosol single-scattering albedo**

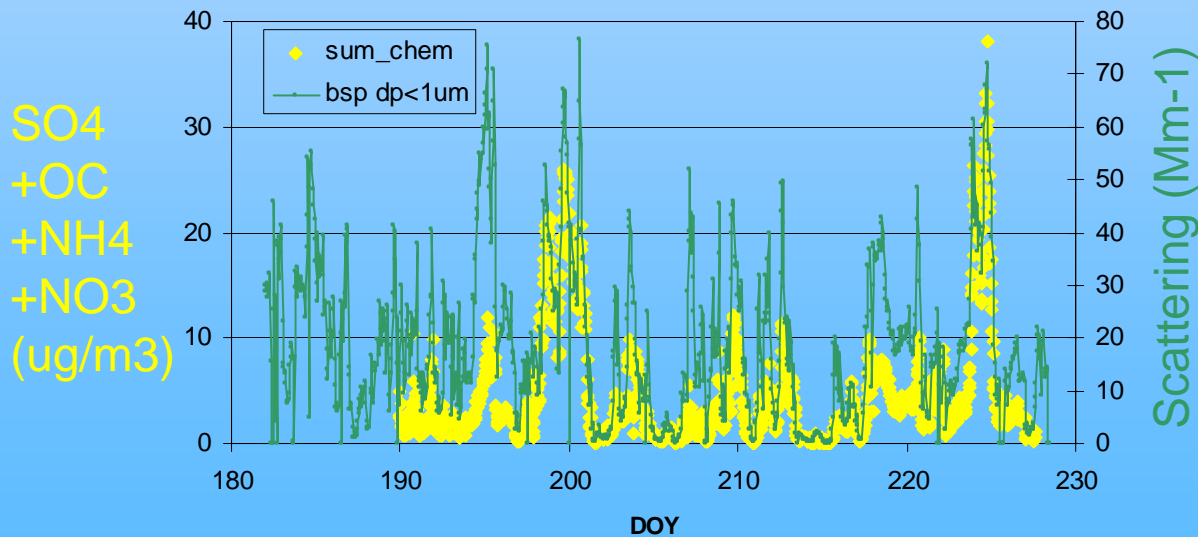
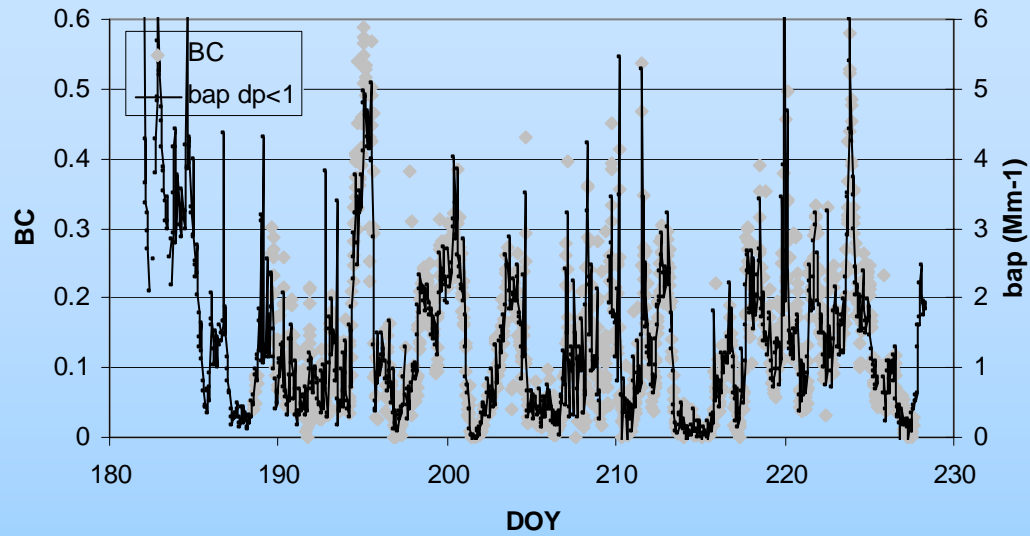


Cloud Scavenging Increases Backscatter Fraction

- **If the particles are externally or internally mixed**
 - light scattering efficiency shows a strong size-dependence in the transition region between scavenged and unscavenged particles
 - light backscattering scattering efficiency shows a weaker size-dependence in the transition region between scavenged and unscavenged particles
 - backscatter fraction shows a very strong size-dependence in the transition region between scavenged and unscavenged particles
- **Cloud droplets are therefore enriched in light scattering particles relative to light backscattering particles**
- **When precipitation falls, it removes more of the light scattering particles than the light backscattering ones**
- **Cloud scavenging therefore systematically increases aerosol hemispheric backscattering fraction**

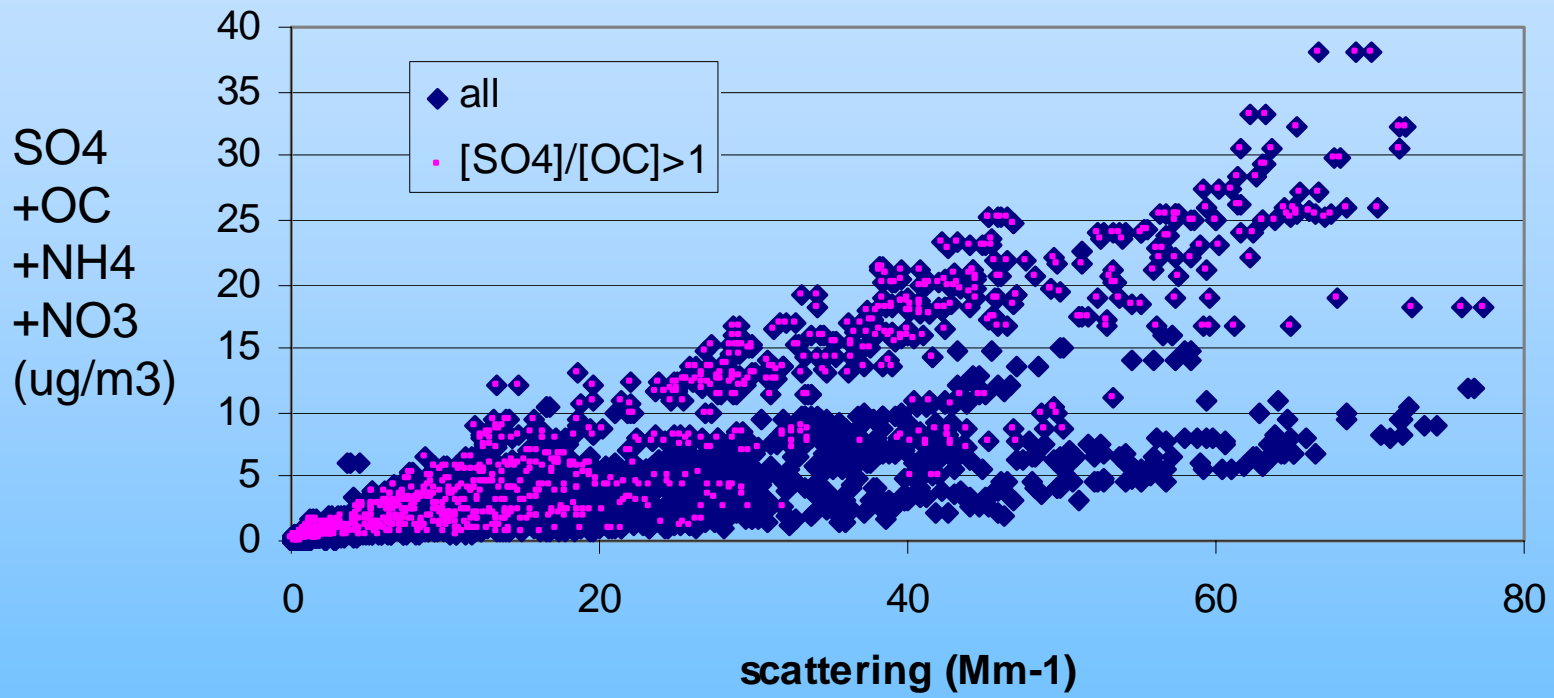


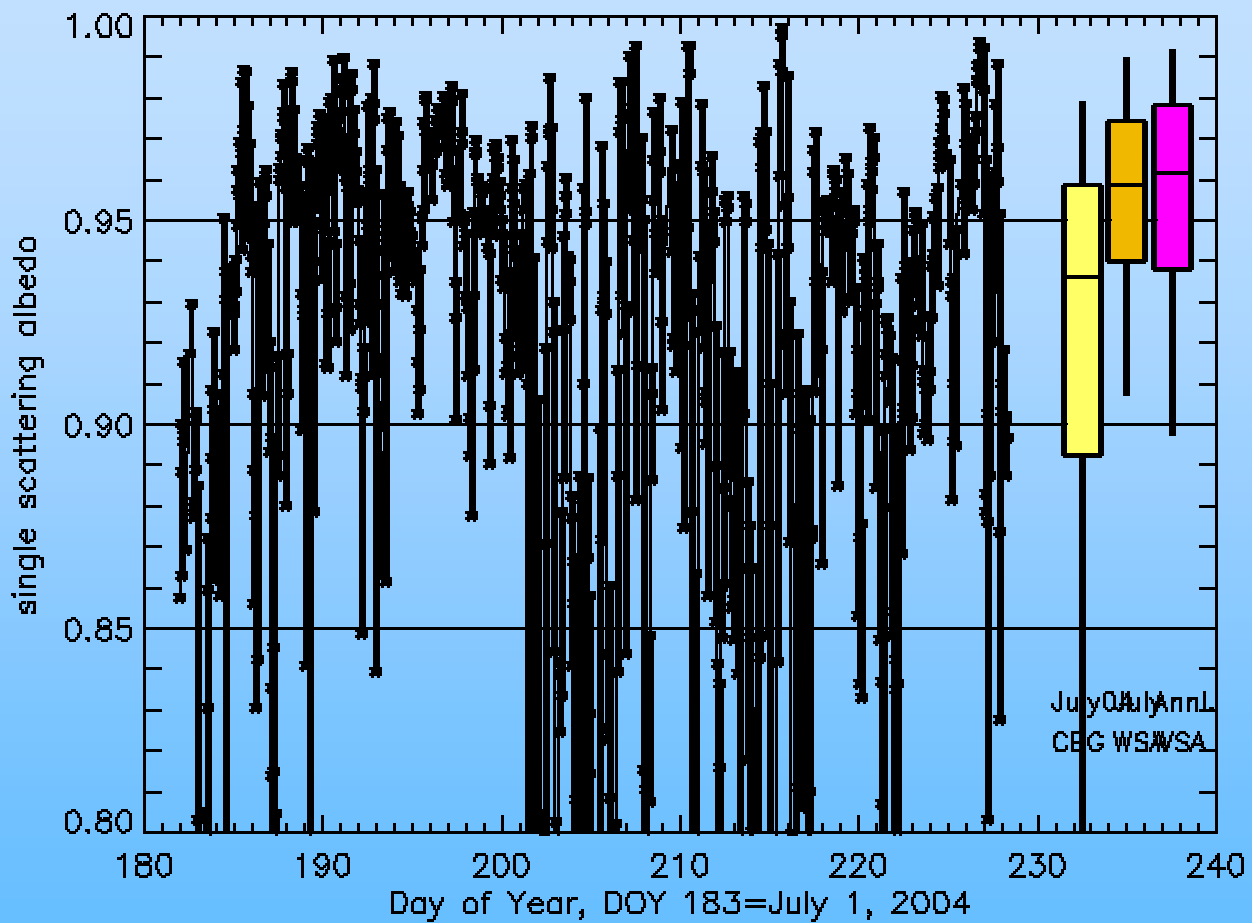
Optical properties and chemistry

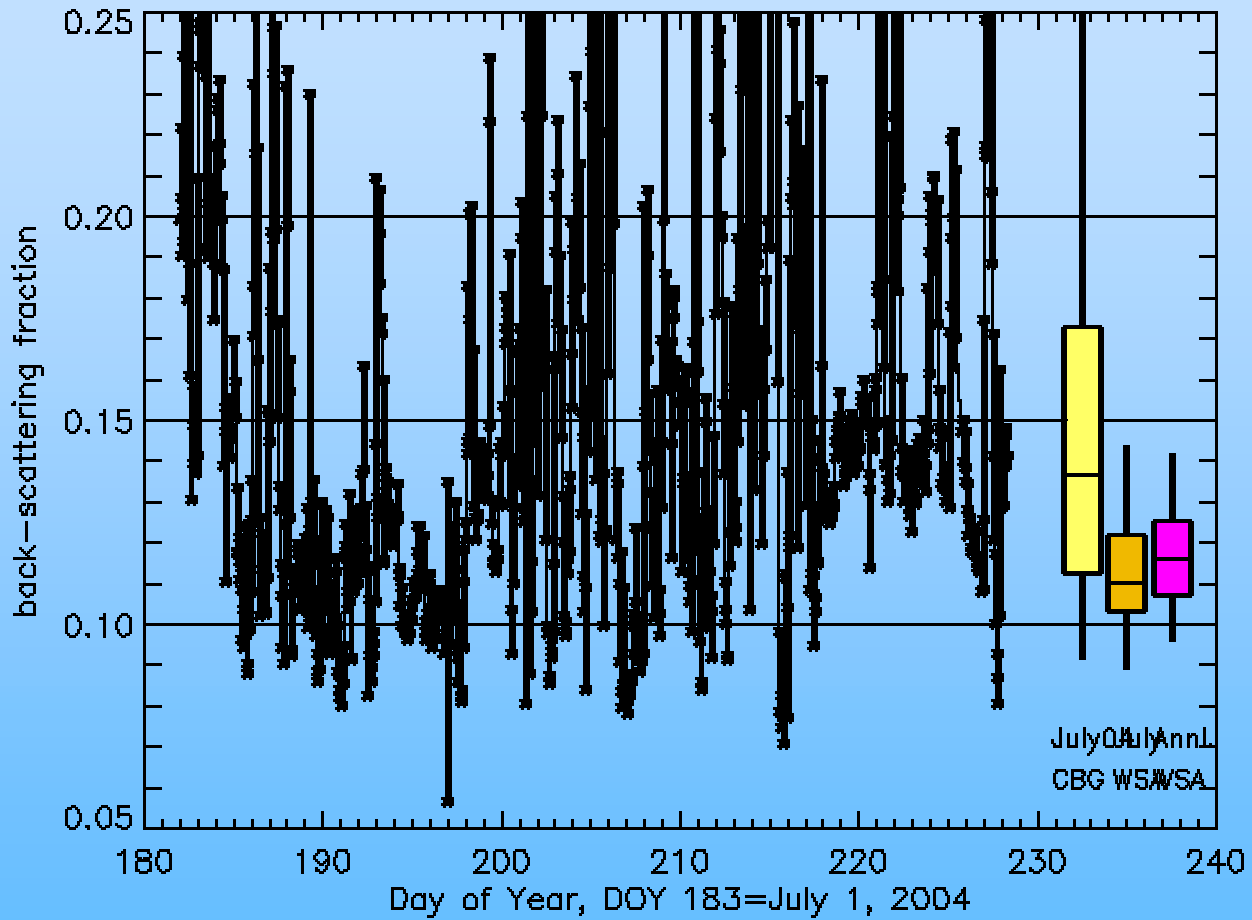


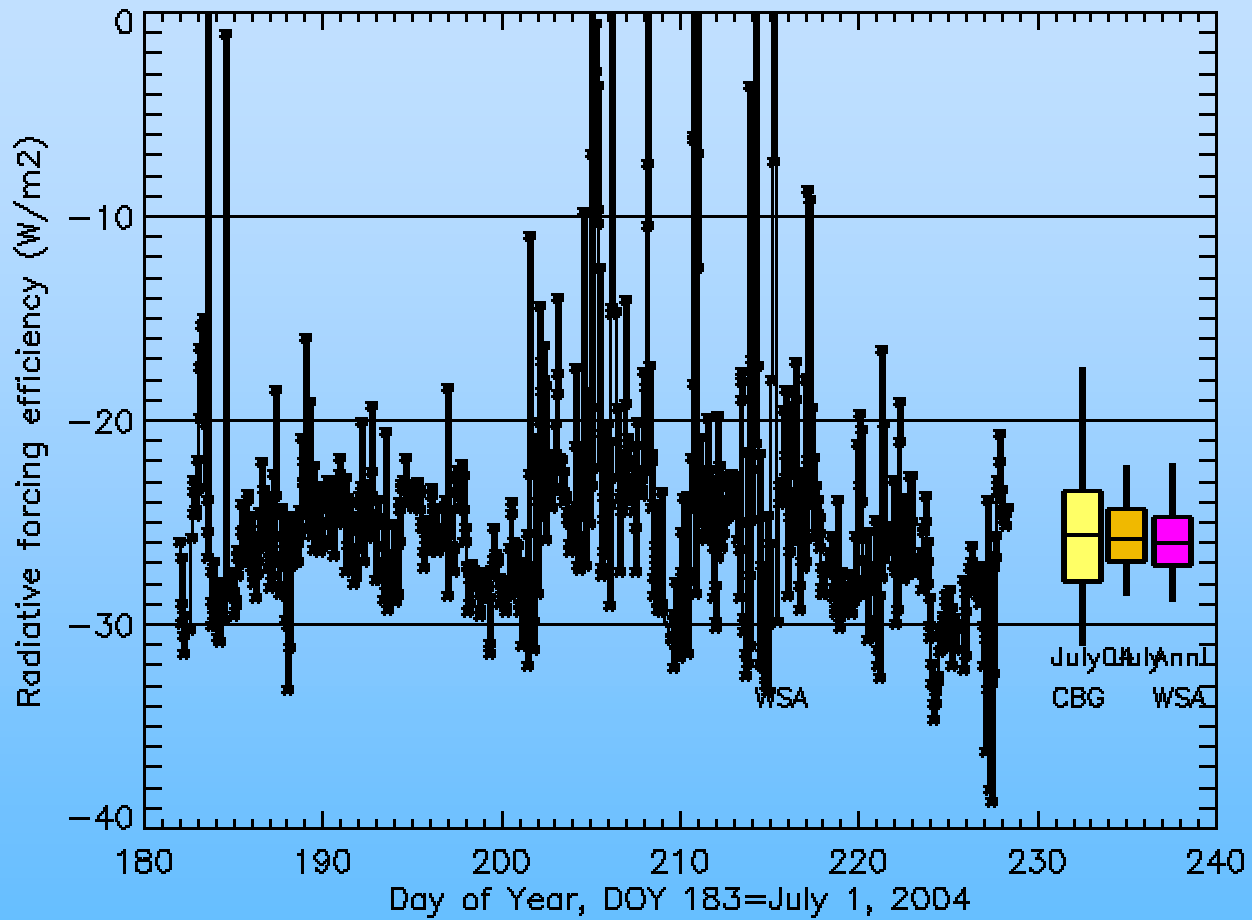
Times when big difference between scattering and sum_chem are when $SO_4 > OC$
→ Other ions? Water?

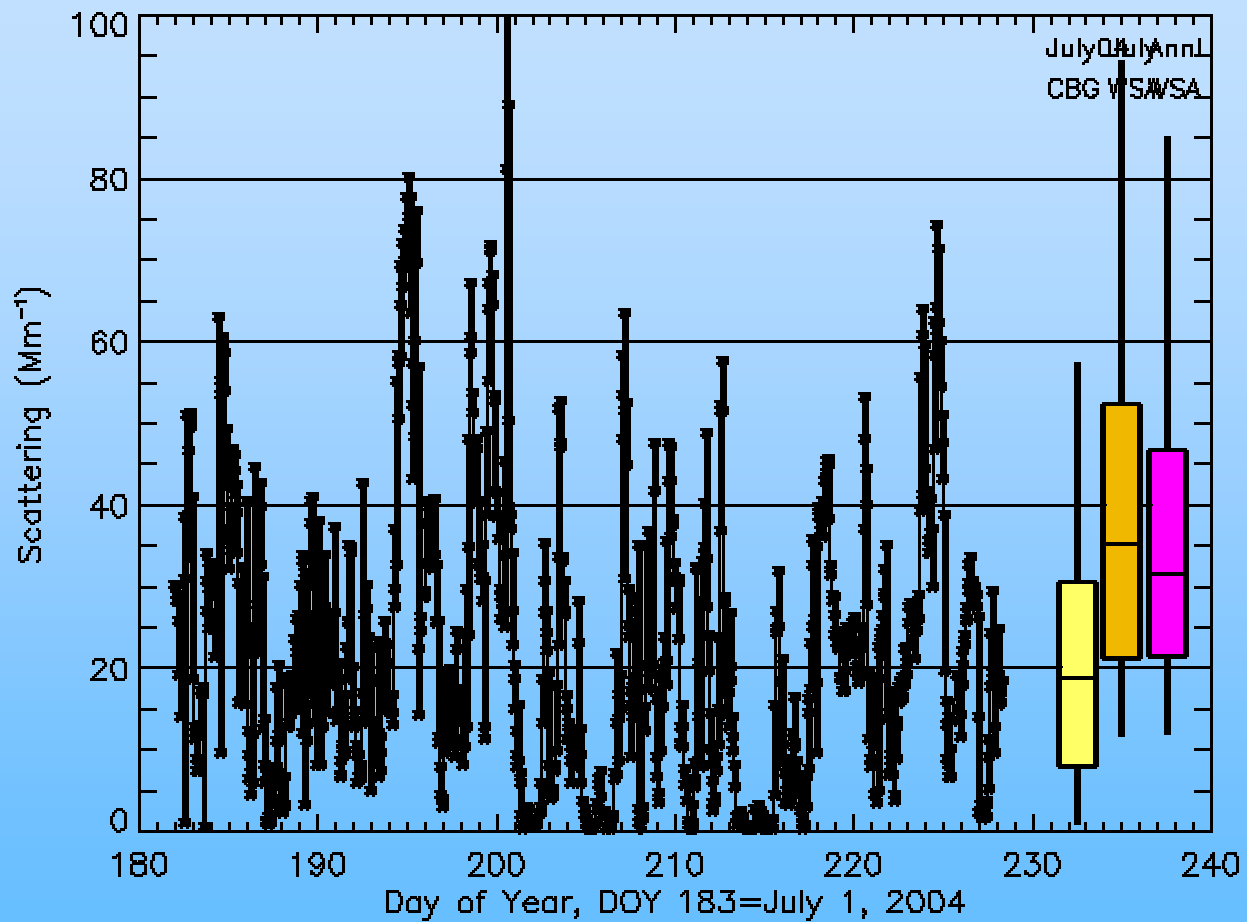












Measurements and derived parameters

MEASUREMENTS

Light scattering (total) (450, 550, 700 nm)

Light scattering (back) (450, 550, 700 nm)

Light absorption (550 nm)

Particle number concentration

Cloud condensation nuclei number concentration at 5 supersaturation values

Ion concentrations and particle mass (PMEL)

Particle size distributions

DERIVED PARAMETERS

Back scatter fraction

Aerosol single scattering albedo

Ångström exponent

