



How Robust is the Observed 11-year Solar Cycle Signal in the Tropical Stratospheric Ozone?

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Introduction

- Understanding the natural decadal variability of stratospheric ozone (O₃) is crucial for monitoring ozone recovery as well as anthropogenic-related changes.
- The currently accepted solar response of the tropical stratospheric O₃ has a **double-peak structure**: one peak of ~3% in the upper stratosphere near 1 hPa and another peak of 3% in the lower stratosphere near 100 hPa. In the broad region of the middle stratosphere (3–30 hPa), the observed response is statistically insignificant (Figure 1).
- The aforementioned solar response has been derived from three long-term satellite instruments: TOMS/SBUV, SAGE I/II, and HALOE.

Unresolved Problems

- From the modeling perspective, the upper stratospheric peak is due to direct photochemical production of O₃ while the lower stratospheric peak is likely an aliasing of volcanic eruptions [Austin et al., 2008; Chiodo et al., 2014].
- However, the mid-stratospheric minimum has not been simulated satisfactorily (Figure 1).
- Proposed mechanisms including QBO and residual circulations and NO_x transport do not fully explain the existence of the minimum [Callis et al., 2001; Dhomse et al., 2011; Kren et al., 2014].

Linear Regression Model

The ozone time series are first deseasonalized to obtain the monthly anomalies. Subsequently, the solar-cycle modulation is retrieved using a regression model delineated in Randel and Cobb [1994] and Li et al. [2008]:

$$X_{O_3}(t) = \alpha(t) \cdot t + \beta(t) \cdot F_{10.7}(t) + \gamma_1(t) \cdot QBO_1(t) + \gamma_2(t) \cdot QBO_2(t) + \varepsilon(t) \cdot ENSO(t) + residual$$

Results and Discussions

- Only the upper stratospheric peak persists in all datasets. But the vertical location of this peak is higher (1 hPa) in the SBUV datasets. In the other datasets, the upper peak is at 3 hPa, which is more consistent with the model simulation.
- The mid-stratospheric minimum derived from the SBUV v8/v8.6 datasets has a broad vertical extent (3–30 hPa), consistent with Soukharev and Hood [2006]. However, in the merged SBUV data, not only the broad mid-stratospheric minimum disappears, but there is also a new third peak in the mid-stratosphere that cannot be explained by models.
- The O₃ solar responses derived from SAGE-II, HALOE, GOZCARDS and SWOOSH do not show a broad minimum in the mid-stratosphere; rather they show a narrower minimum at lower altitudes that better agrees with the model simulations.

Aim

- Re-examine the tropical stratospheric O₃ solar response using long-term O₃ data publicly available online.

Data

- SBUV MOD v8.0
- SBUV MOD v8.6
- SBUV Merged Cohesive (CDR) 1979 onwards
- SAGE-II 1984–2005
- HALOE 1991–2005
- GOZCARDS Merged SAGE-I/II, HALOE, ACE-FTS, UARS-MLS, Aura-MLS 1979–2012
- SWOOSH Similar to GOZCARDS except without ACE-FTS and different merging methods 1984–2012

Conclusions

- The mid-stratospheric minimum of the tropical stratospheric O₃ solar response only appears in SBUV v8/v8.6 datasets.
- Cross-instrument validation is required to evaluate the O₃ data quality on decadal time scales.

References

Austin et al (2008), JGR, **113**, D11306; Soukharev & Hood (2006), JGR, **111**, D20314; Li et al (2008), JGR, **113**, D14109; Randel & Cobb (1994), JGR, **99**, 5433; Dhomse et al (2011), **11**, 12773; Callis et al (2001), **106**, 7523; Chiodo et al (2014), ACP, **14**, 5251.

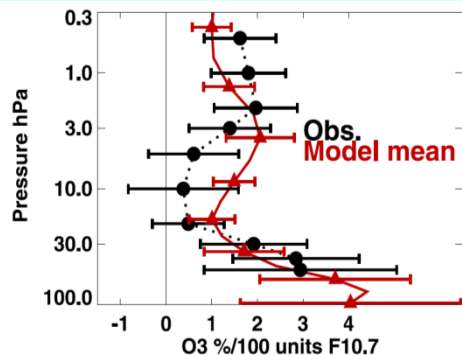


Figure 1. O₃ solar-cycle response averaged over the latitude range 25°S to 25°N. The red line shows simple mean of simulations of seven chemistry-climate models with solar forcing. The dotted black line is the mean of the observations from three independent satellite instruments presented by Soukharev & Hood [2006]. The uncertainty ranges are 95% confidence intervals. Extracted from Austin et al. [2008].

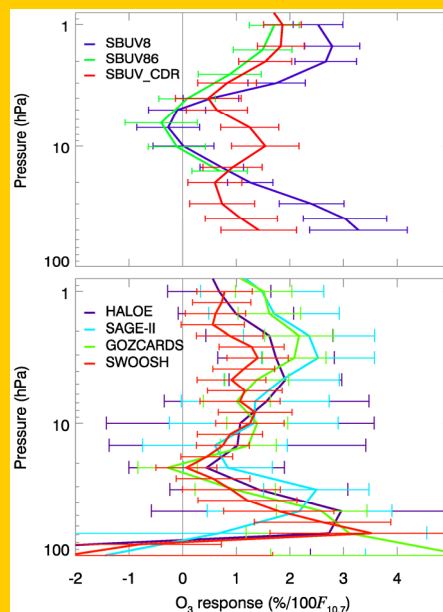


Figure 2. The vertical solar responses of tropical stratospheric O₃ derived from various satellite datasets.